UNIVERSITÄT LEIPZIG

Climate Dynamics (Summer Semester 2017)

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Today's Lecture (Lecture 5): General circulation of the atmosphere

Reference

- Hartmann, *Global Physical Climatology* (1994), Ch. 2, 3, 6
- Peixoto and Oort, Ch. 4, 6, 7

2.3 – General circulation of the atmosphere

- Atmospheric transport in response to radiative imbalance
- Mean meridional circulation and eddy circulation

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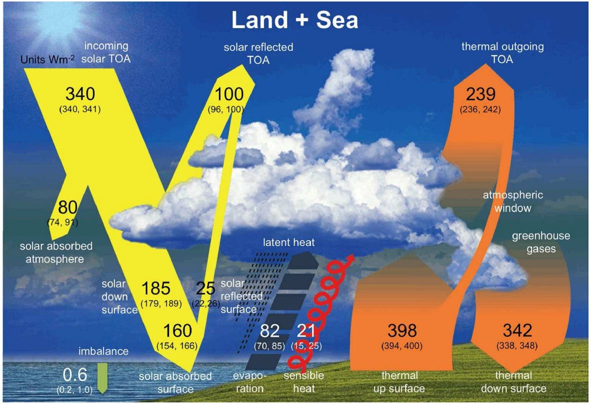
- Mean meridional circulation and eddy circulation
- Energy cycle
- Entropy cycle
- Cycles of momentum, angular momentum
- Hydrological cycle

Radiative budgets for the atmosphere and at TOA



Radiative energy balance of the atmosphere (sign convention: downwelling positive) is

$$
\begin{aligned}
R_a = F_{\mathrm{TOA}} - F_s + R_{\mathrm{TOA}} - R_s &= (340 - 100) - 160 + (-239) - (342 - 398)\ \mathrm{W\ m^{-2}} \\
&= \mathcal{O}(-100\ \mathrm{W\ m^{-2}}),
\end{aligned} \tag{2.85}
$$

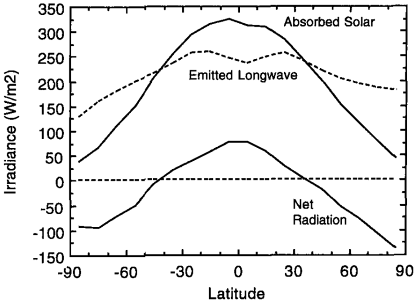
balanced by fluxes of sensible and latent heat into the atmosphere

Figure: Wild et al., 2015

Zonal-mean radiative budget at TOA

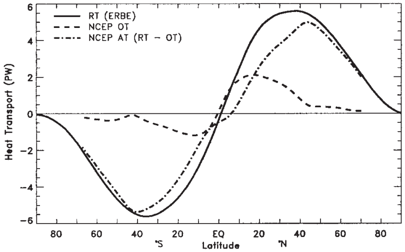
TOA radiative budget. . .

- The top-of-atmosphere (TOA) radiative balance measures how much energy enters or leaves the climate system
- In the tropics, the net energy flux is positive (into the climate system)
- In the extratropics, the net energy flux is negative (out of the climate system)



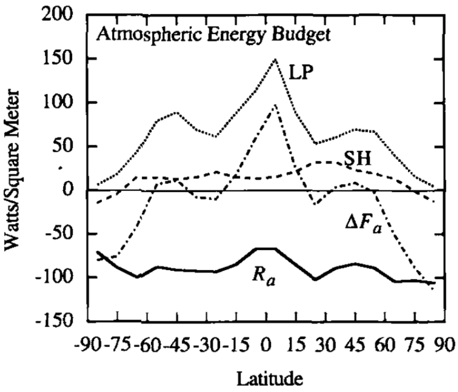
. . . requires meridional energy transport

- To maintain a steady state (e.g., constant long-term zonal-mean temperature), energy transport from the tropics into the extratropics is required
- The *meridional divergence* of the energy transport balances the radiative energy flux
- Contributions to transport from ocean and atmosphere



Figures: Hartmann (1994), Trenberth and Caron (2001)

Atmospheric energy budget requires atmospheric transport



- As we saw in the previous section, the net radiative energy balance of the atmosphere is $R_a = \mathcal{O}(-100 \text{ W m}^{-2})$; the balance is fairly constant in latitude
- The radiative energy loss is balanced by latent (LP) and sensible (SH) heat flux from land and ocean; but these are strong functions of latitude
- Meridional advective atmospheric energy flux is required to provide local energy balance:

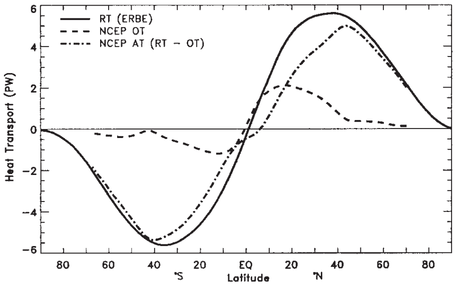
$$R_a + F_{\text{LH}} + F_{\text{SH}} = \Delta F_a \tag{2.86}$$

- The advective energy flux is the meridional divergence of the meridional heat transport (sign convention: northward positive):

$$\frac{dN}{d\phi} = \int_0^{2\pi} d\lambda\, R_E^2\, \Delta F_a(\phi) \cos\phi = 2\pi R_E^2\, \Delta F_a(\phi) \cos\phi \tag{2.87}$$

Figures from Hartmann (1994) unless noted

Atmospheric energy budget requires atmospheric transport



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Trenberth and Caron (2001)

Averaging operators

Temporal mean

$$\overline{A} = \overline{A}(\lambda, \phi, p) = \frac{1}{\tau} \int_{-\tau/2}^{\tau/2} A(\lambda, \phi, p, t)\, dt \tag{2.88}$$

and the *zonal mean*

$$[A] = [A](\phi, p, t) = \frac{1}{2\pi} \int_{0}^{2\pi} A(\lambda, \phi, p, t)\, d\lambda \tag{2.89}$$

The instantaneous value of A is given by

$$A = \overline{A} + A' \tag{2.90}$$

where A' is called the *fluctuating* component of A. Likewise

$$A = [A] + A^* \tag{2.91}$$

where A^* is the departure from the zonal mean.
Decomposition of a field into time-average and fluctuating, zonally symmetric and zonally asymmetric components:

$$A = [\bar{A}] + [A'] + \bar{A}^* + A'^* \tag{2.92}$$

Decomposition of the flow

Products of fields contain covariance terms (where fluctuations do not average to zero)

$$\overline{AB} = \bar{A}\bar{B} + \overline{A'B'} \tag{2.93}$$

$$[AB] = [A][B] + [A^*B^*] \tag{2.94}$$

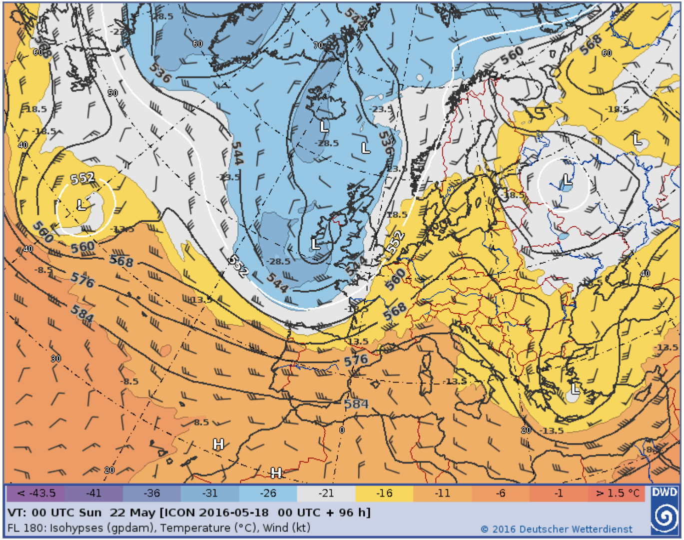
$$\left[\overline{AB}\right] = \left[\bar{A}\right]\left[\bar{B}\right] + \left[\bar{A}^*\bar{B}^*\right] + \left[\overline{A'B'}\right] \tag{2.95}$$

The terms in (2.95) are the mean circulation, stationary eddies, and transient eddies. To take a concrete example, the decomposition of northward flux of sensible heat is

$$c_p\left[\overline{vT}\right] = c_p\left[\bar{v}\right]\left[\bar{T}\right] + c_p\left[\bar{v}^*\bar{T}^*\right] + c_p\left[\overline{v'T'}\right] \tag{2.96}$$

This week's homework will analyze the relative importance of each contribution as a function of latitude.

Example of meridional heat transport by a transient eddy



$[v][T] \ll [v^* T^*]$ in the eddy covering northern Europe

Mean circulation – streamfunction

The zonal-mean continuity equation (zonal flow is integrated out) is

$$\frac{1}{R_E \cos\phi} \frac{\partial}{\partial \phi}([\bar{v}] \cos\phi) + \frac{\partial [\bar{\omega}]}{\partial p} = 0 \tag{2.97}$$

For a nondivergent flow, velocity components can be written with the aid of a streamfunction:

$$[\bar{v}] = \frac{g}{2\pi R_E \cos\phi} \frac{\partial \Psi_M}{\partial p} \tag{2.98}$$

$$[\bar{\omega}] = \frac{-g}{2\pi R_E^2 \cos\phi} \frac{\partial \Psi_M}{\partial \phi} \tag{2.99}$$

(2.98) and (2.99) satisfy (2.97); normalization, including the minus sign, is convention – but the relative minus sign is required. To calculate Ψ_M, first impose boundary condition $\Psi_M = 0$ at TOA, then integrate (2.98):

$$\Psi_M = \frac{2\pi R_E \cos\phi}{g} \int_0^p [\bar{v}]\, dp' \tag{2.100}$$

The normalization is chosen to give units of kg s^{-1} (*mass streamfunction*); the $\cos\phi$ factor is required to ensure constant Ψ_M for constant meridional flow. Mass transport is tangent to contours of the streamfunction. Mass flow between two contours is equal to $\Delta\Psi_M$.

Mean meridional circulation

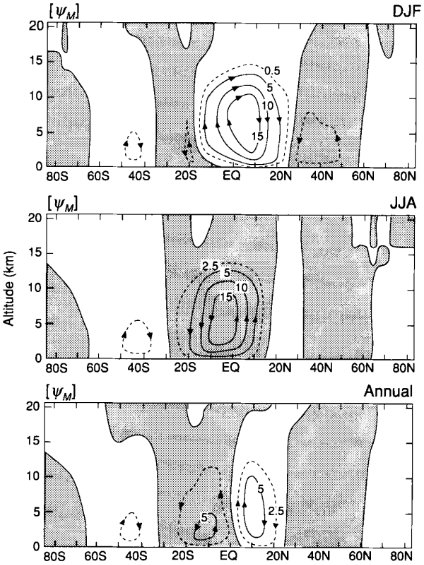
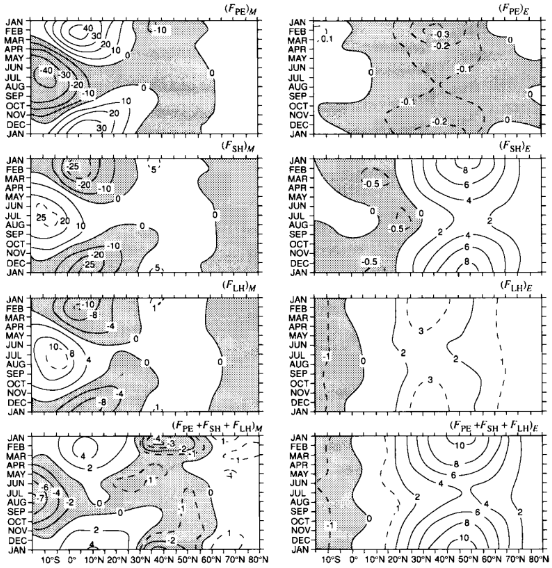


Figure: Hartmann (1994); shaded: $\Psi_M < 0$; units: 10^{10} kg s^{-1}

- Hadley cell with rising branch in the ITCZ and descending in the subtropics
- Transport is from winter hemisphere to summer hemisphere at the surface, summer hemisphere to winter hemisphere at altitude → transport of potential energy, latent heat, sensible heat
- Mass transport by mean circulation is small outside the Hadley cell
- This is where (temporal and zonal) fluctuations in the circulation are important – eddy transport

Meridional energy transport



- Recall static energy (2.72): sum of potential energy (PE), sensible heat (SH) and latent heat (LH)

$$h = gz + c_p T + L_{lv} q \qquad (2.101)$$

  The divergence of poleward transport of these energy terms balances the atmospheric energy budget.

- Mean transport dominates in the Hadley cell – but note large terms of opposite signs
- Eddy transport, especially in winter (large temperature gradient), dominates in midlatitudes