Today's Lecture: Land, biosphere, cryosphere (All that stuff we don't have equations for...)

Reference

- UNEP *Global Outlook for Ice and Snow* (2007)
- IPCC AR5
- NSIDC
- (all linked from course web page)

4 – Land, biosphere, cryosphere

Reference

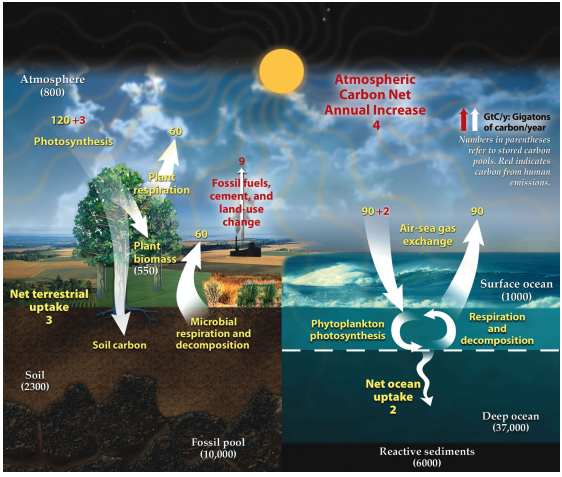
UNEP *Global Outlook for Ice and Snow* (2007), IPCC AR5, NSIDC (all linked from course web page)

4.1 – Land–biosphere–atmosphere–ocean interaction

- Land is a sink of atmospheric momentum
- Orography shapes circulation (stationary Rossby waves)
- Land–sea temperature contrast shapes circulation
- Land and ocean are a source/sink of sensible and latent heat
- Land and ocean are a source of aerosol
- Land and ocean are a source/sink of trace gases

4.2 – The carbon cycle

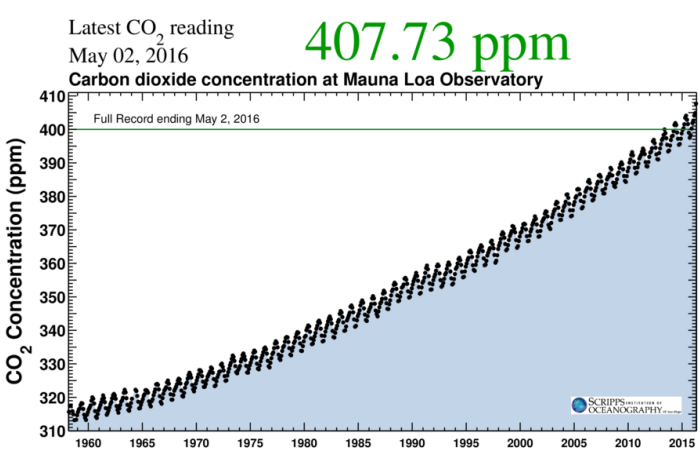
Carbon cycle and carbon reservoirs

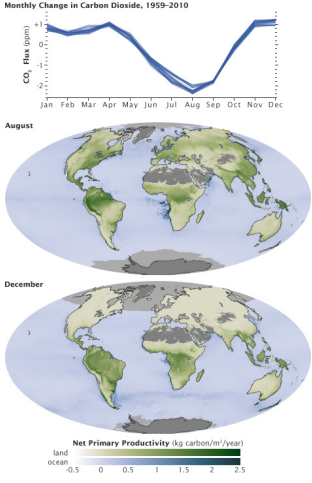


- Carbon reservoirs are large, but cycling is slow
- Anthropogenic carbon fluxes are small compared to the natural fluxes
- But the flux imbalance is large compared to the natural flux imbalance
- Only about 50% of emitted anthropogenic carbon remains in the atmosphere in the short term

Figure: U.S. Department of Energy Office of Science

The fast carbon cycle – seasonal cycle of biological primary productivity





Figures: Scripps CO_2 program, NASA, NOAA ESRL

The slow carbon cycle: weathering, biogeochemical pump, metamorphism, volcanoes

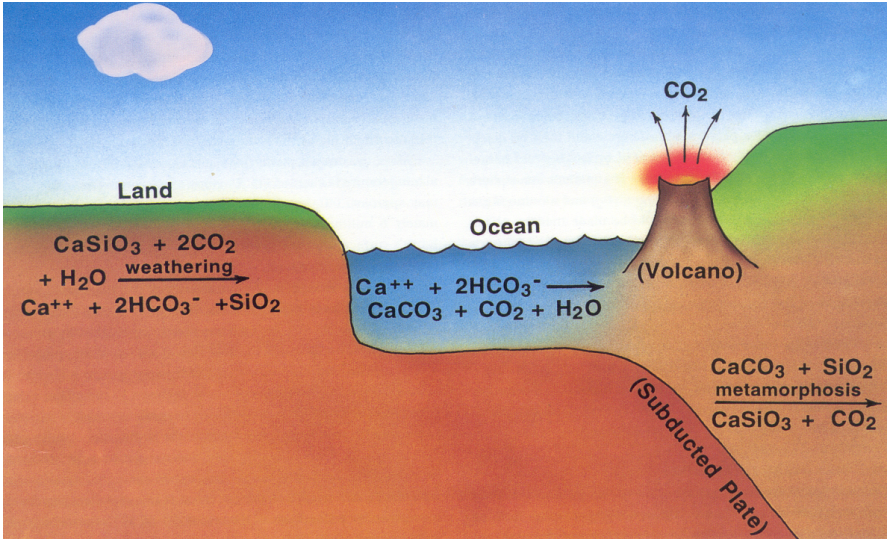
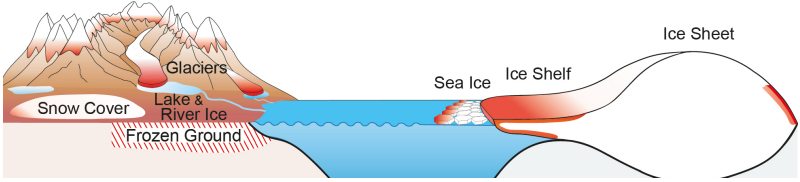


Figure: Kasting 1995

4.3 – Cryosphere–atmosphere–ocean interaction

- The cryosphere acts as a reservoir for water, which is released on short (annual) and long (> millennial) time scales
- Freezing and melting are strong local influences on ocean salinity
- Albedo of ice affects shortwave flux into ocean
- Low thermal conductivity insulates ocean from atmosphere

Components of the cryosphere



- Sea ice
- Ice sheets
- Ice shelves
- Glaciers

Which of these (directly) influence the sea level?

Figure: IPCC AR5

4.4 – Sea ice

- Annual cycle of freezing and melting
- First-year and multi-year ice; ice thickness, persistence through melt season
- Ice albedo (depends on snow cover)
- *Polynyas* as source of sensible and latent heat

Annual cycle of sea ice extent

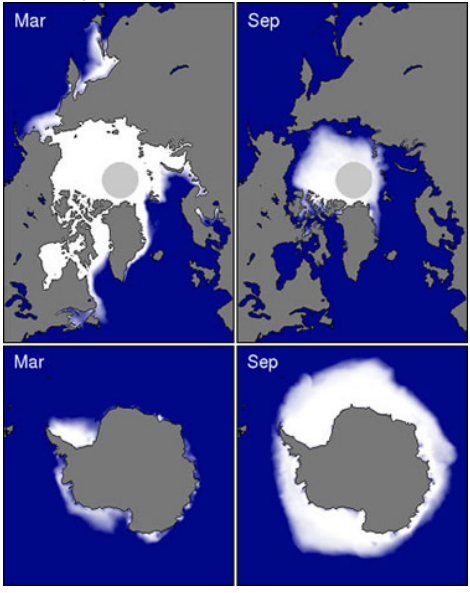
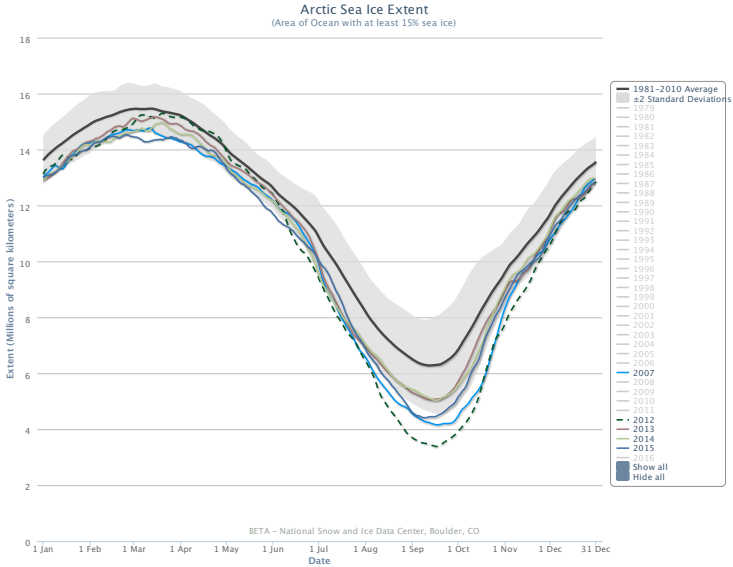


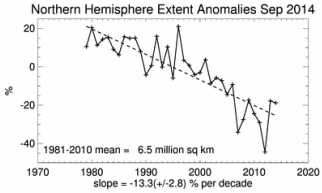
Figure: NSIDC

- Arctic sea ice occupies the Arctic Ocean, including the pole; partly persists for multiple years
- Antarctic sea ice forms equatorward of the Antarctic continent and consists mostly of first-year ice
- Freezing of the Arctic Ocean restricts moisture flux $\rightarrow$ Arctic sea ice is polar desert with low snow cover (bare ice albedo: 0.5)
- Southern Ocean provides moisture source for snowfall on Antarctic sea ice (snow-covered ice albedo: 0.9)

Annual cycle of arctic sea ice extent

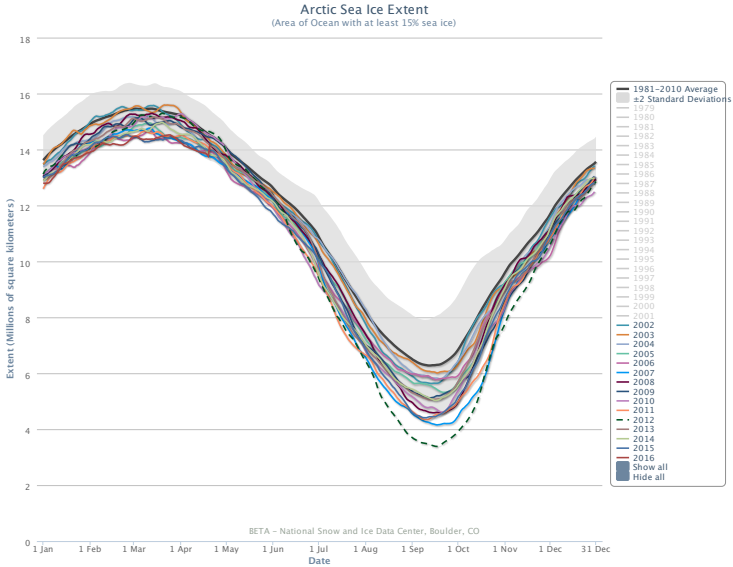


- Annual cycle is much larger than interannual variability
- Interannual variability is also large compared to the trend
- The trend is very large compared to zero (anomaly sign is the same year after year)



Figures: NSIDC

Annual cycle of arctic sea ice extent



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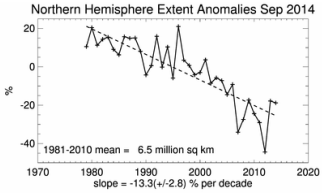
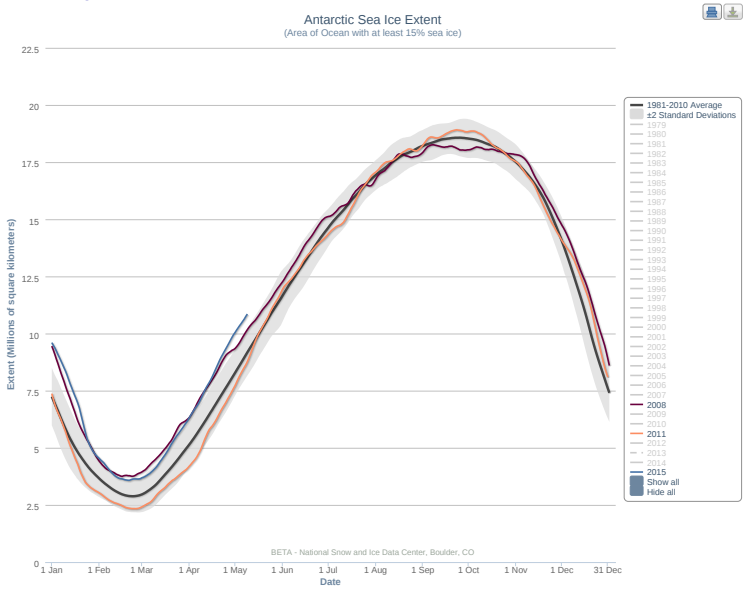
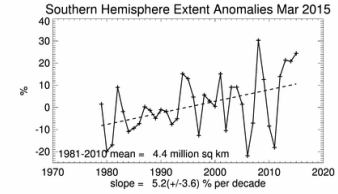


Figures: NSIDC

Annual cycle of antarctic sea ice extent



- Annual cycle is much larger than interannual variability
- Interannual variability is also large compared to the trend
- The trend is small (anomaly sign is often different between years) and positive



Figures: NSIDC

Ice thickness

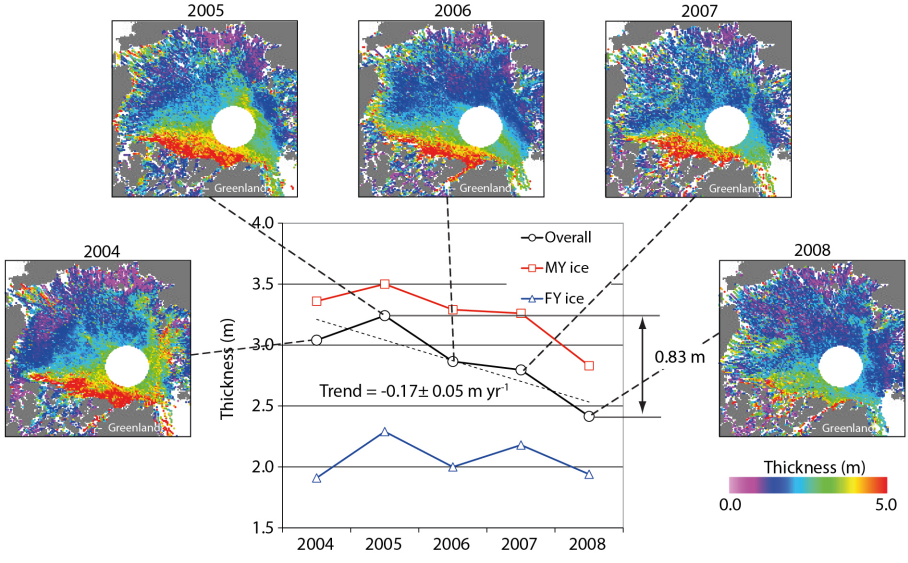


Figure: IPCC AR5

4.5 – Continental ice sheets and ice shelves

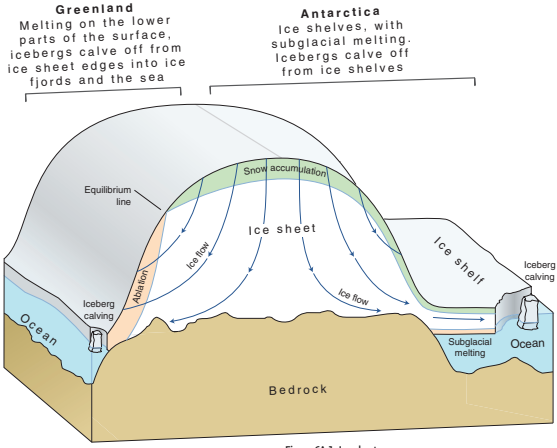


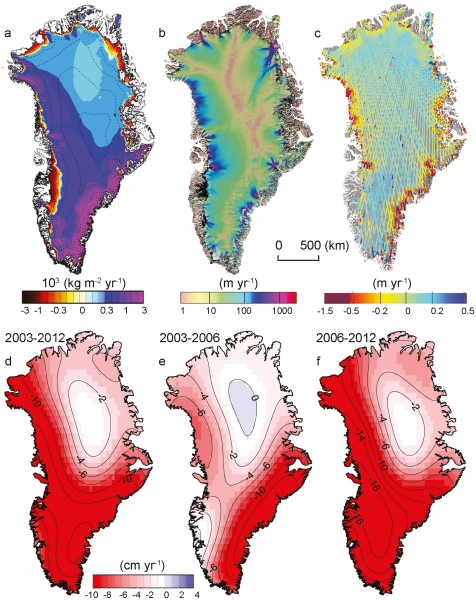
Figure 6A.1: Ice sheets.

Source: based on material provided by K. Steffen, CIRES/Univ. of Colorado

- *Ice sheets* are accumulations of permanent (i.e., non-seasonal) ice of continental size
- In the present-day climate, there are two: Greenland and Antarctica
- Whether their mass increases or decreases (the *mass balance*) depends on snow accumulation rate (mass source) and melting and iceberg calving (mass sinks).
- Depending on temperature, warming can result in mass gain (due to increased snow fall) or mass loss (melting, faster ice flow, reduced back pressure from collapsed ice shelves)
- Equivalent sea level rise is 60 m (Antarctica) and 7 m (Greenland); crucial to know whether, when, and how much of the ice sheets will melt
- Dynamics depend on basal lubrication (difficult of observe), but satellite gravimetry and altimetry provide the flow field (since ca 2000, with gaps)

Figure: UNEP *Global Outlook for Ice and Snow* (2007)

Observed mass balance of Greenland



a Model-derived accumulation

b Flow speed (satellite)

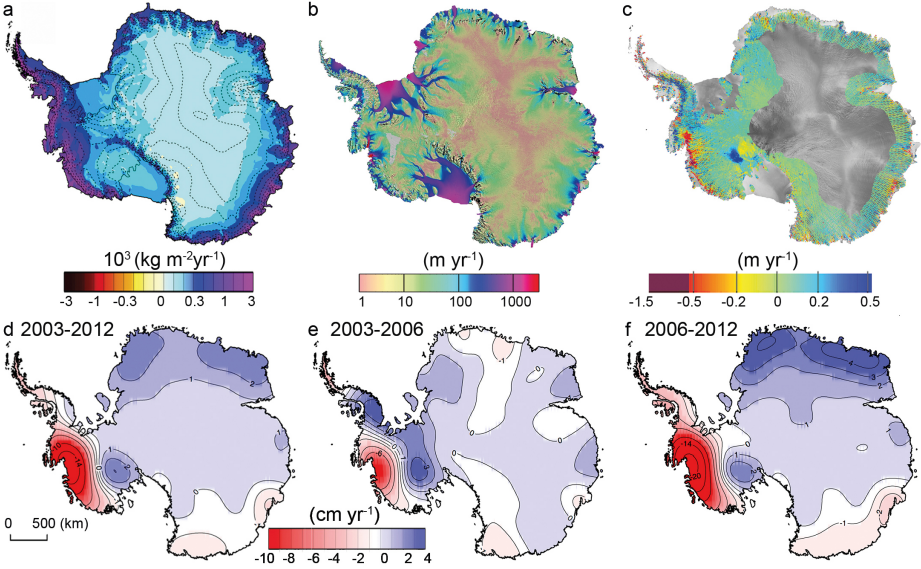
c Elevation change (satellite)

d Ice loss (cm water yr^{-1}, gravimetry), 2003–2012

e Ice loss (cm water yr^{-1}, gravimetry), 2003–2006

f Ice loss (cm water yr^{-1}, gravimetry), 2006–2012

Observed mass balance of Antarctica



Figures: IPCC AR5

Ice shelf collapse (Larsen B, 2002)

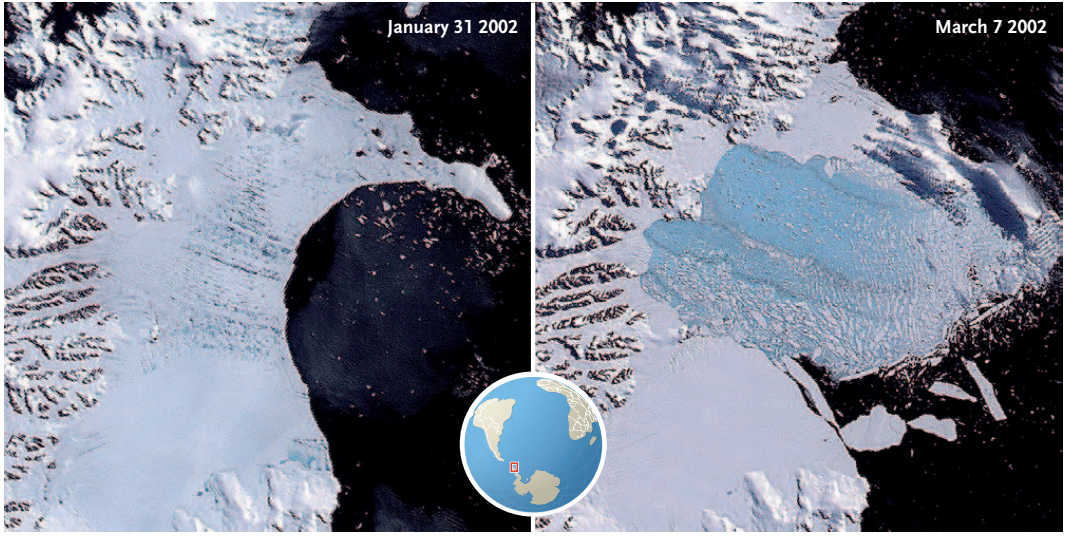
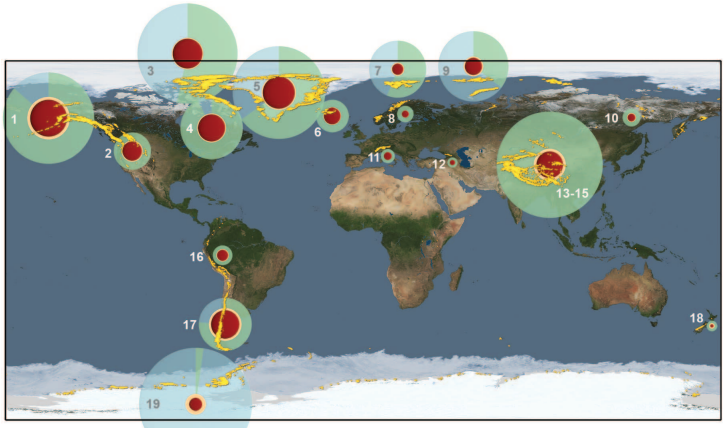
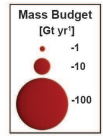
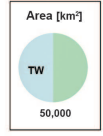


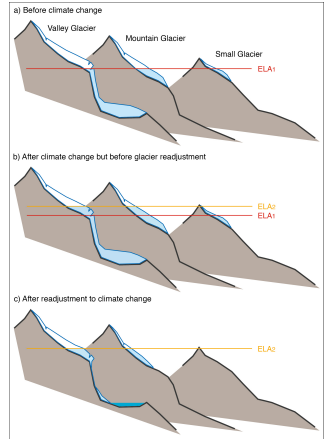
Figure: UNEP *Global Outlook for Ice and Snow* (2007)

4.6 – Glaciers









- Sea level equivalent is small (< 1 m)
- But the are an important water source in tropics and subtropics
- Universally in decline, with very few exceptions
- Glacier response lags warming, so further decline is *committed*

Figures: Gardner et al. (2008), IPCC (2013)